

Modelling the fluid and heat flows below the TVZ: the challenges of fluid properties

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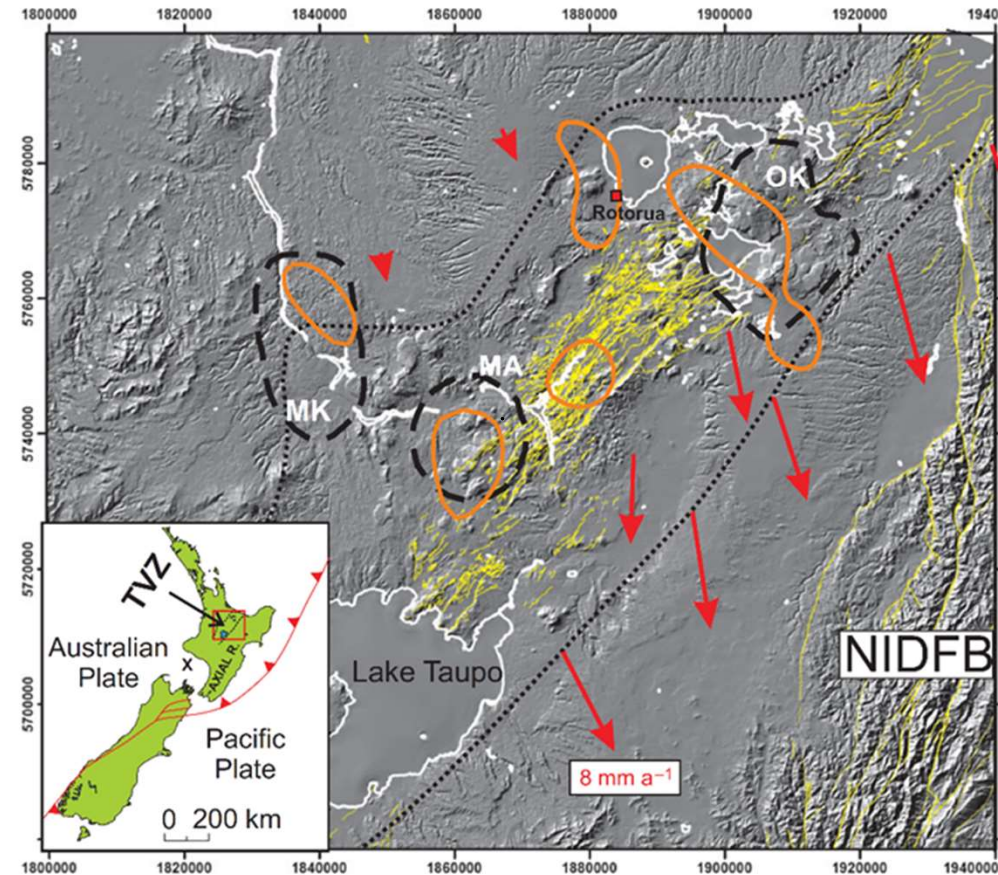
GNS Science

Taupo Volcanic Zone (TVZ)

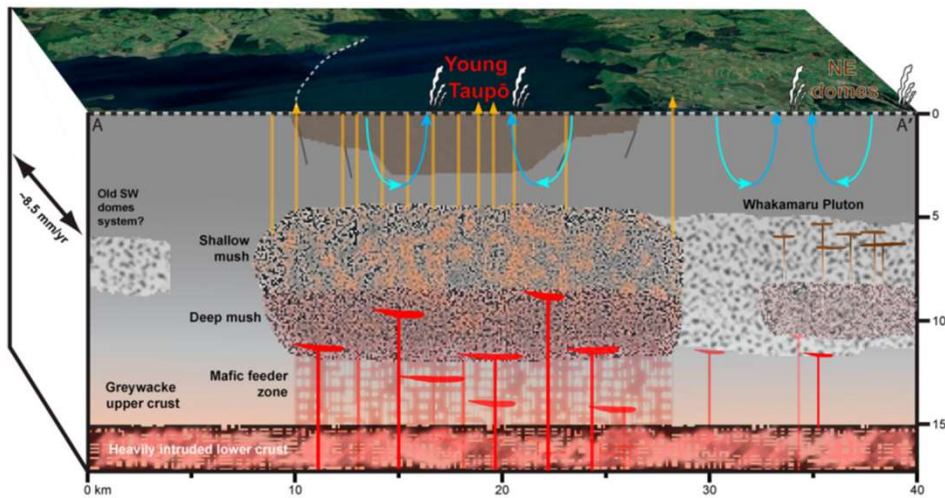
Regional-scale fluid and heat flow modelling with realistic deep heat source ($\sim 900^{\circ}\text{C}$, 0.7 W/m^2)

How does the heat get from $\sim 15 \text{ km}$ deep to near-surface at $100\text{-}300^{\circ}\text{C}$?

What do super-critical geothermal resources look like ?



Super-critical geothermal resources



Barker et al., NZJGG, 2020.

Shallow magma ‘mush’ or a magma intrusion is likely to underlie super-critical resources in the TVZ

Temperatures of $\sim 900^{\circ}\text{C}$ possible at < 5 km depth

Vigorous shallow convection and boiling is challenging to model

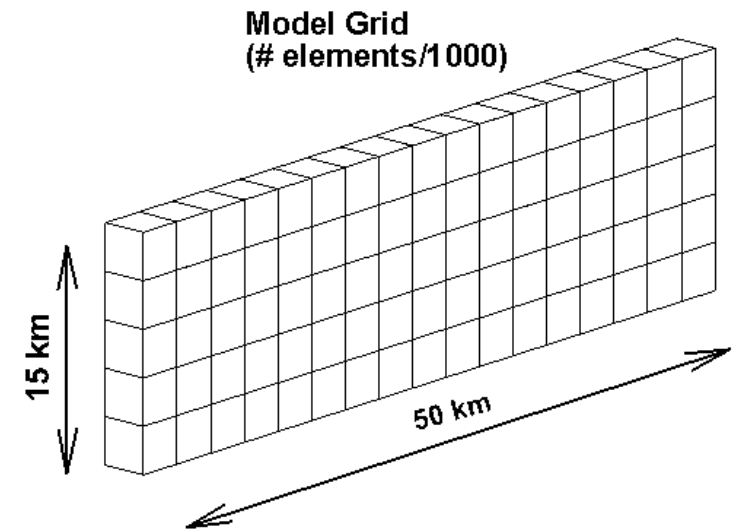
Fluid and heat flow modelling for geothermal

TGNS is an in-house fluid and heat transport research code

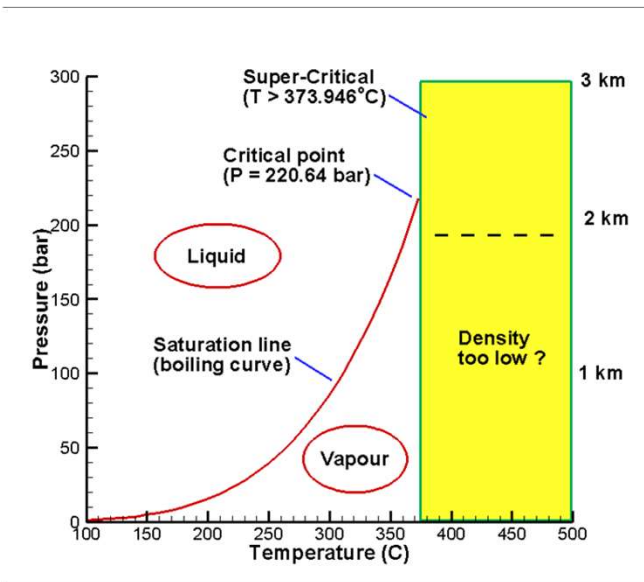
Similar 'fully implicit' algorithm to Tough2 but with extended thermodynamics

Solves the equations describing how water and heat move together in the Earth's crust

Uses IAPWS-95 and can handle pressures and 'magmatic' temperatures at base of crust



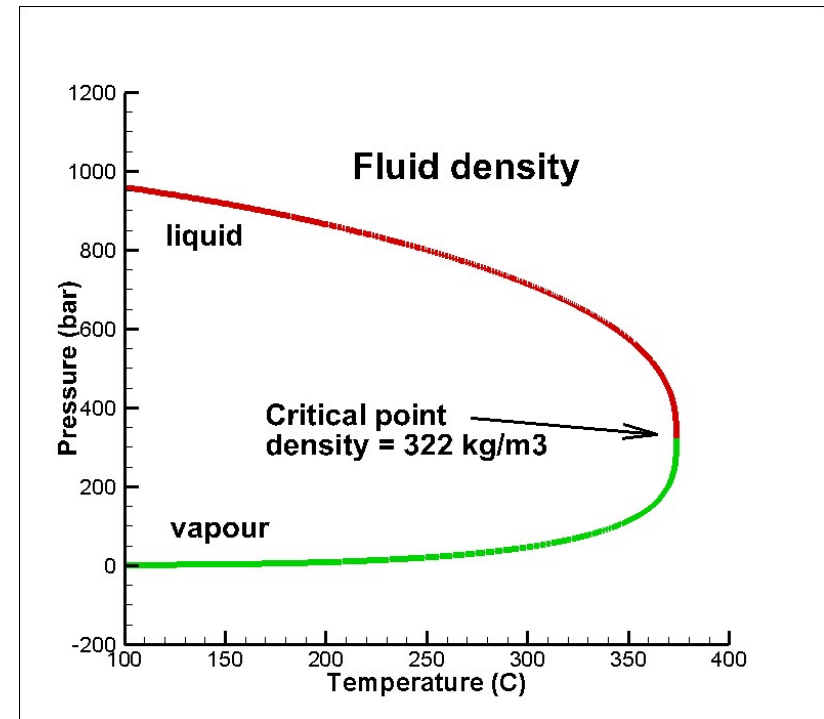
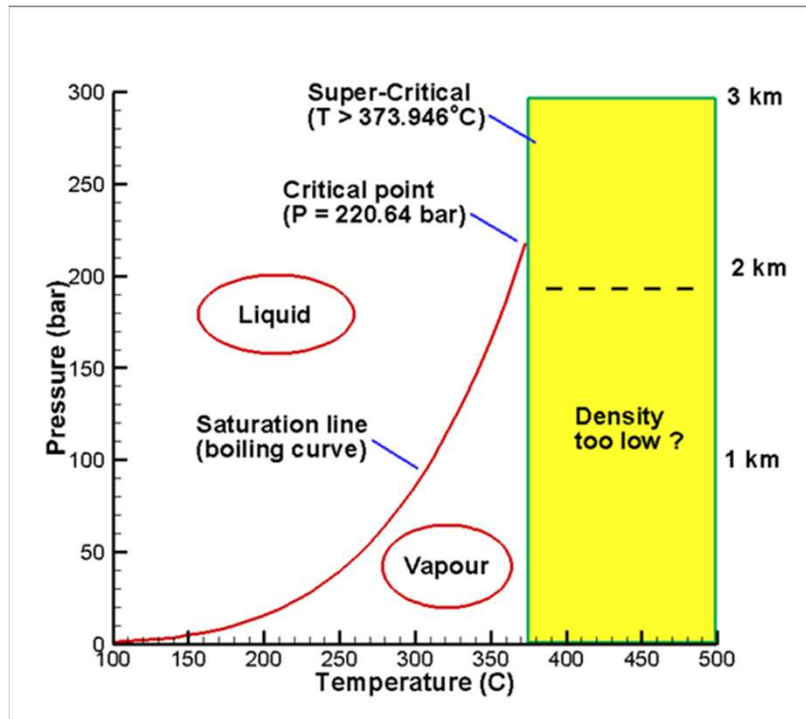
Properties of (liquid) water



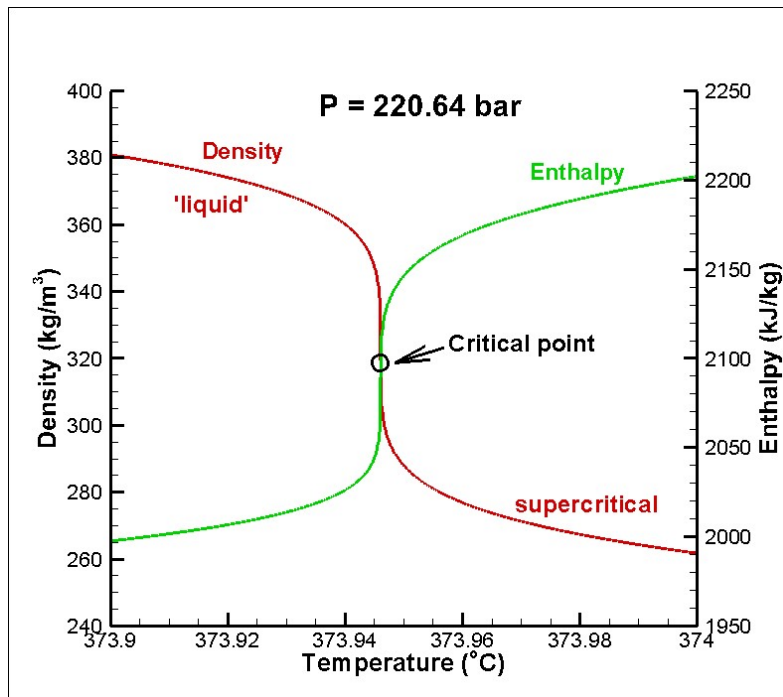
| | Atmospheric conditions | 'Normal' geothermal | Super-Crit. geothermal | Base of crust (15km) |
|----------------|---------------------------|----------------------|------------------------|----------------------|
| | 1 bar, 20°C | 100 bar, 300°C | 300 bar, 380°C | 1000 bar, 900°C |
| Density | 998 kgm ⁻³ | 715 | 534 | 198 |
| Enthalpy | 84 kJkg ⁻¹ | 1343 | 1838 | 4056 |
| Dyn. viscosity | 1 x 10 ⁻³ Pa.s | 9 x 10 ⁻⁵ | 6 x 10 ⁻⁵ | 5 x 10 ⁻⁵ |

Differences between hot and cold fluid properties make for challenging modelling

Saturation line and liquid/vapour phase properties



Phase properties near the critical point

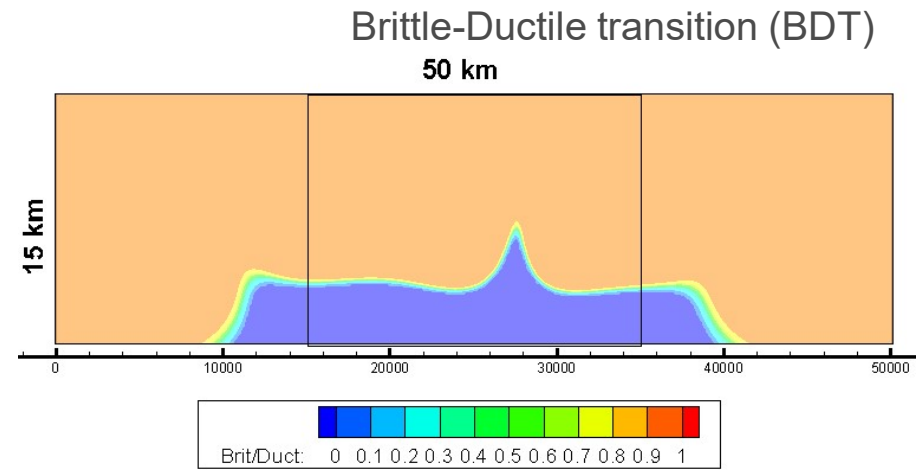
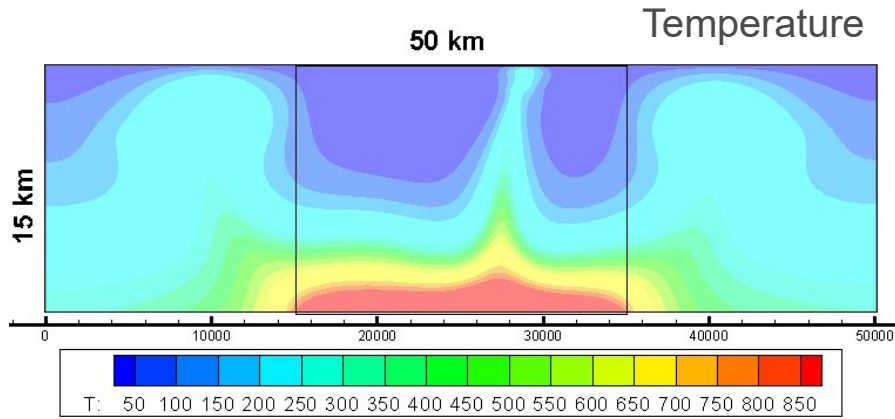


Sub-critical: single- and two-phase fluids behave 'normally'

Near-critical point: two-phase fluids very different from 'normal' and steep gradients difficult to deal with numerically

Better high-temperature phase changing algorithm remains elusive

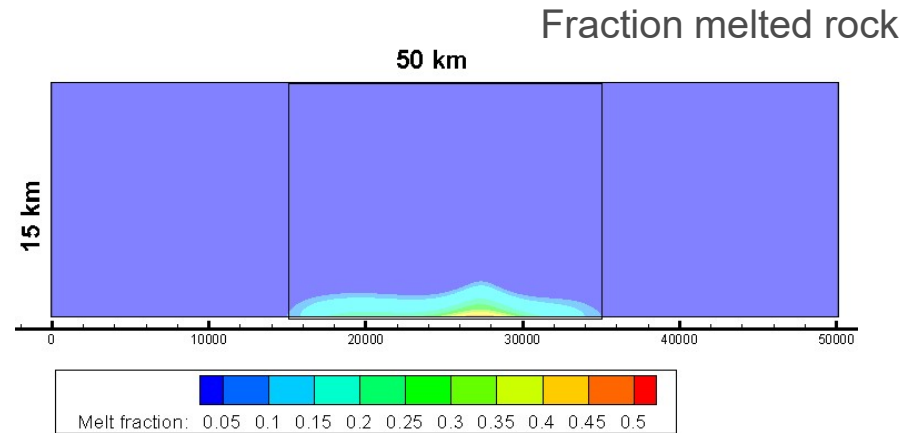
Examples of model results



No boiling at surface and the 'geothermal systems' are too cool

B-D transition is too deep

Too much melted rock



Conclusions: IAPWS-95 formulation

IAPWS-95 formulation uses density and temperature as independent variables

Pressure is more 'intuitive' than density. Requires finding pressure corresponding to density and temperature (iterative, expensive and prone to convergence problems)

Water properties must be 'mathematically' smooth and continuous functions of pressure (density) and temperature
IAPWS-95 has largely fixed the smoothness problems in earlier formulations, except (possibly) on the saturation line near critical point

Conclusions: modelling/algorithms

High-temperature boiling (say $> 330^{\circ}\text{C}$) is difficult to achieve
This occurs above *any* magma/mush body and so it is important to know how to handle it...

In the near-super-critical point region the gradients of thermal properties of water with respect to P & T are large and cause spurious results and lack of convergence
Maybe avoid implicit solution methods ??